Chapter 2: Heat and Water Vapor Transfer from Vegetation

Waring RH and SW Running. 1998. <u>Forest Ecosystems, Analysis at Multiple</u> <u>Scales, 2nd ed</u>. Academic Press, USA. Ch.2, Water cycle, 22-28.

Estimating environmental parameters

- Equations are based on the idea of energy balance
- Use values you can measure directly or estimate easily to estimate hard to determine processes
- Our goal is to figure out what happens to the water present in the system

A. Energy Balance

- Movement of water vapor is determined using:
 - Radiant energy in the system (Net Radiation)
 - Environmental variables controlling evaporation
 - Temperature of the environment compared to the surface
 - Humidity of the environment compared to the air very close to the surface
 - Movement of air

Net Radiation

- Net radiation is comprised of:
 - Short wave radiation 95% of solar radiation is short wave
 - Long wave radiation
- Total radiation striking a surface comes from:
 - Direct short wave- sun
 - Diffuse short wave- sky
 - Reflected short wave- nearby surfaces
 - Long wave atmospheric emission
 - Long wave emission from nearby surfaces

R_n =(1- α) I_s + $ε_L \sigma T^4$ (surface) - σT^4 (sky) 2.1 Net Radiation = short wave radiation + long wave radiation



Surface Energy Balance

- Net radiant energy on a surface is comprised of:
 - Energy stored by the surface as heat (G)
 - Normally this is a small value on short timescales
 - Sensible heat flux (H) Heat changing the air temperature
 - Value is hard to measure
 - Latent heat flux (λE) Heat used to evaporate water

$R_n - G = H + \lambda E \qquad 2.2$ Net Radiation – Heat Storage = Sensible Heat Flux + Latent Heat Flux



Bowen Ratio (β)

- Relationship between sensible and latent heat losses. (H/ λ E)
- Used to determine Latent Heat when Sensible Heat is unknown.
- Need to estimate Bowen ratio based on known information for this to work.

 $\lambda E = R_n / (1 + \beta) 2.3$

B. Evaporation from Wet Surfaces

- Movement of water vapor is determined using:
 - Radiant energy in the system (Net Radiation)
 - Environmental variables controlling evaporation
 - Temperature of the environment compared to the surface
 - Humidity of the environment compared to the air very close to the surface
 - Movement of air

Simple Evaporation Rate $E = (e_w - e_a) f(u)$ 2.4

Evaporation rate = evaporation from the surface * air circulation

Evaporation from the surface $(e_w - e_a)$ is determined by the vapor pressure of the air at the surface compared to the surrounding air

- $e_w > e_a$ Water evaporates from the surface into the air
- $e_w = e_a$ Local equilibrium so no evaporation net change in water
- e_w < e_a Water condenses from the air to the surface

Increasing turbulence will increase Evaporation.



Finding Evaporation without Surface Temperature

• Using the Penman equation:

$$E = E_{eq} + D g_b / \zeta_{2.5}$$

In this equation

• E_{eq} is the evaporation from a specific surface at equillibrium

 $E_{eq} = (\epsilon / \rho \lambda (\epsilon + 1)) R_n \qquad 2.6$

 ζ (zeta) is a combination of constants and easily measured variables

 $\zeta = \rho (\epsilon + 1) G_v T_k$

$E = E_{eq} + Dg_b / \zeta \quad 2.5$

- E_{eq} Evaporation rate at equilibrium
- D Air saturation deficit
 - Warm dry air near the surface will increase this value increasing evaporation
- g_b Boundary layer conductance
 - Small values (low conductance) will make surrounding air conditions less important and make evaporation more similar to equilibrium conditions



Pictures from:

http://www.meted.ucar.edu/nwp/model_physics/print.htm